



The Catchment Isoscape: A Meta-Model for Stable Isotope Tracers at the Shale Hills Critical Zone Observatory

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In this paper:

- Describe a 5 year experiment for water isotopes at Shale Hills CZO
- Present a theory for "age" and "residence time" of mobile-immobile water flow in soils and regolith
- Compare theory and experiment to test for the existence of macroporematrix flow at Shale Hills

"The Catchment Isoscape"



Advancing interdisciplinary studies of earth surface processes

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The Susquehanna/Shale Hills Critical Zone Observatory





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Shale Hills Observations & Experiments to Support Earth System Models & Prediction











Stable Isotope Experiment 2008-20012 Holmes, Thomas, Gaines, Jin, Andrews, Lin, Duffy $\delta^{2}H - \delta^{18}O$ Stream



Soil Moisture - Xylem Water Experiment





Catchment $\delta^2 H$ Signature From Time Series



A Single Dominant Period: 360 < T < 370 days



Seasonal Cycle Attenuates over Profile



The Catchment Isoscape & The Age of Water

The term "isoscape" was coined by ecologists to describe the spatial and temporal patterns of isotope ratios over a landscape (Bowen, 2010)

In this paper we explore an "isoscape" for water isotopes of $\delta^2 H - \delta^{18} O$ in soil and regolith

The goal is to develop a meta-model for spacetime isotopic patterns over the catchment as a step towards predicting:

"The Distributed Age and Residence Time of Water Isotopes at the Shale Hills CZO"

Properties of the Age Distribution

$$\mu_n(x,t) = \int_0^\infty \tau^n c(x,\tau,t) d\tau$$

$$\mu_o(x,t) \Rightarrow C(x,t) = \int_0^\infty c(x,t,\tau) d\tau$$

$$\mu_1(x,t) \Rightarrow \alpha(x,t) = \int_0^{\infty} \tau c(x,t,\tau) d\tau$$

$$A(x,t) = \frac{\mu_1(x,t)}{\mu_o(x,t)} = Mean Age of water$$

A Model for Age Distribution

Rotenberg 1972, J, of Theoretical Biology, 37, 291-305





$$\mathbf{D}M(t,\tau)\frac{1}{V} = \left(\frac{\partial M}{\partial t} + \frac{\partial M}{\partial \tau}\right)\frac{1}{V}$$

$$\frac{\partial c}{\partial t} + \frac{\partial c}{\partial \tau} = \Gamma_c + L(c)$$

$$L(c) \rightarrow D \frac{\partial^2 c}{\partial x^2} - u \frac{\partial c}{\partial x}$$

or

$$L(c) \Longrightarrow \frac{Q_i}{V}(c_i - c)$$

Transport Model in Terms of Moments

$$\frac{\partial c}{\partial t} + \frac{\partial c}{\partial \tau} = \Gamma_c + L(c)$$



Source terms for the nth moment

A Theory for Concentration-Age For Mobile-Immobile Transport Over the Soil Profile







$$\frac{d\alpha_{im}}{dt} = C_{im} + \frac{k}{\theta_{im}}(\alpha_m - \alpha_{im})$$

$$A_{m}(t) = \alpha_{m}(t) / C_{m}(t);$$

$$C_{m}(0,t) = C_{i}(t); C'_{m}(d,t) = 0;$$

$$C_{m}(z,0) = C_{im}(z,0) = 0;$$

$$A_{im}(t) = \alpha_{im}(t) / C_{im}(t)$$
$$\alpha_m(0,t) = \alpha_{im}(0,t) = \alpha'_m(d,t) = 0$$
$$\alpha_m(z,0) = \alpha_{im}(z,0) = 0$$



Simulated Seasonal Input C_i (scaled)







Simulated Isotope Ratio





Simulated Age of





Modeling the Catchment Isoscape for Shale Hills

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G. Bhatt 2012



Penn State Integrated Hydrologic Model (PIHM)



Semi-Discrete Approach: PIHM

Process	Governing equation/model	Original governing equations	Semi-discrete form
Channel Flow	St. Venant	$\frac{\partial h}{\partial t} + \frac{\partial (uh)}{\partial x} = q$	$\left(\frac{d\varsigma}{dt} = P_c - \sum Q_{gc} + \sum Q_{oc} + Q_{in} - Q_{out} - E_c\right)_i^{[1]}$
Overland Flow	Equation	$\frac{\partial h}{\partial t} + \frac{\partial (uh)}{\partial x} + \frac{\partial (vh)}{\partial y} = q$	$\left(\frac{\partial h}{\partial t} = P_o - I - E_o - Q_\infty + \sum_{j=1}^3 Q_j^{(j)}\right)_i^{[1]}$
Unsaturated Flow	Richard Equation	$C(\psi)\frac{\partial\psi}{\partial t} = \nabla\cdot(K(\psi)\nabla(\psi+Z)$	$\left(\frac{d\xi}{dt} = I - q^0 - ET_s\right)_i^{[2]}$
Groundwater Flow		$C(\psi)\frac{\partial\psi}{\partial t} = \nabla\cdot(K(\psi)\nabla(\psi+Z)$	$\left(\frac{d\zeta}{dt} = q^{0} + \sum_{j=1}^{3} Q_{g}^{ij} - Q_{l} + Q_{gc}\right)_{i}^{[3]}$
Interception	Bucket Model	$\frac{dS_I}{dt} = P - E_I - P_o$	$\left(\frac{dS_I}{dt} = P - E_I - P_o\right)_I$
Snowmelt	Temperature Index Model	$\frac{dS_{snow}}{dt} = P - E_{snow} - \Delta w$	$\left(\frac{dS_{snow}}{dt} = P - E_{snow} - \Delta w\right)_{t}$
Evapotranspiration	Pennman- Monteith Method	$ET_{0} = \frac{\Delta(R_{n} - G) + \rho_{a}C_{p} \frac{(e_{s} - e_{a})}{r_{a}}}{\Delta + \gamma(1 + \frac{r_{s}}{r_{a}})}$	$\left(ET_0 = \frac{\Delta(R_n - G) + \rho_a C_p \frac{(e_r - e_a)}{r_a}}{\Delta + \gamma(1 + \frac{r_r}{r_a})}\right)_r$

Distributed IsotopeTransport (Bhatt, 2012)

Transport Equation

$$\theta A_B \left[C_i \frac{\partial h_i}{\partial t} + h_i \frac{\partial C_i}{\partial t} \right] = - \oint_{\partial \Omega i} \mathbf{n} \cdot (\mathbf{V}\bar{C}) \, dA + \oint_{\partial \Omega i} \mathbf{n} \cdot (\mathbf{D}\nabla\bar{C}) \, dA + \int_{\Omega i} q_S C_S \, d\Psi$$

Semi-Discrete Form

$$\frac{\partial C_i}{\partial t} = \sum_{\partial \Omega i} \frac{\bar{Q}}{\theta A_B h_i} [C_i - \bar{C}] + \sum_{\Omega i} \frac{Q_S}{\theta A_B h_i} [C_i - \bar{C}_S] + \sum_{\partial \Omega i} \frac{\mathbf{n} \cdot (\mathbf{D} \nabla \bar{C}) A_B}{\theta A_B h_i}$$

Horizontal Advection

Vertical Flux

Dispersive Flux

CZO Data ->lidar, Soil, Regolith, Veg



Simulated Average Age of Groundwater + Soil Water At Shale Hills CZO

G. Bhatt, PhD 2012



Random Input C_i and Q_i





Relative Frequency Groundwater Age & Runoff Residence Time

G. Bhatt, PhD 2012



1979-2010 Reanalysis Forcing & Dynamic Residence Time of Runoff



Atmospheric Modeling of Stable Isotopes $\delta^{18}O$ in Precipitation A new research product

IsoRSM experiment over northeast US

10km Simulation covering 85.5W-71.3W/35.5N-46.2N

Boundary Conditions: IsoGSM simulation based on NOAA Climate Reanalysis

Kei Yoshimura, University of Tokyo, Japan (Yoshimura et al., 2010)



Validation of IsoRSM Stable Isotopes in Precipitation with Shale Hills CZO Data



Evan Thomas, MS Penn State

Conclusions:



The "Mean Age" of waters can be simulated directly using transport theory and stable isotopes

The particular form of the age- or transit-time distribution function is not necessary in this theory.

Mobile-Immobile storage attenuates the seasonal amplitude through lateral diffusive exchange, increasing the relative age of infiltrating waters

Vegetation using immobile water is not detached from the mobile phase but rather exchange occurs by capillary diffusion

The isoscape is a powerful concept for assessing space-time patterns of age and residence time at the catchment scale

Thank You